

Record-Setting Ocean Warmth Continued in 2019

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Human-emitted greenhouse gases (GHGs) have resulted in a long-term and unequivocal warming of the planet (IPCC, 2019). More than 90% of the excess heat is stored within the world's oceans, where it accumulates and causes increases in ocean temperature (Rhein et al., 2013; Abram et al., 2019). Because the oceans are the main repository of the Earth's energy imbalance, measuring ocean heat content (OHC) is one of the best way to quantify the rate of global warming (Trenberth et al., 2016; Von Schuckmann et al., 2016; Cheng et al., 2018). Following reports released in the previous two years (Cheng and Zhu, 2018; Cheng et al., 2019c), this article presents new OHC data for the year 2019. These data reveal that the world's oceans (especially at upper 2000 m) in 2019 were the warmest in recorded human history. Specifically, the ocean heat anomaly (0–2000 m) in 2019 was 228 Zetta Joules (ZJ, 1 ZJ=10²¹ Joules) above the 1981–2010 average and 25 ZJ above 2018 (Table 1).

The OHC values (for the upper 2000 m) were obtained from the Institute of Atmospheric Physics (IAP) ocean analysis (see “Data and methods” section, below), which uses a relatively new method to treat data sparseness and updates in the instruments that have been used to measure ocean temperature (Cheng et al., 2017). The evolution of OHC (Fig. 1) shows that the upper 2000 m OHC in 2019 was 228 ± 9 ZJ above the 1981–2010 average. The record-setting ocean warmth is also found in National Oceanic and Atmospheric Administration/National Center for Environmental Information (NOAA/NCEI) data, showing 217 ± 4 ZJ in 2019 above 1981–2010 average (21 ZJ above 2018) (Table 1) (updated from Levitus et al. 2012). With these newly available IAP data, a ranking of the warmest years since 1950s is now possible (Table 1). The past five years are the top five warmest years in the ocean historically with modern instruments, and the past ten years are also the top ten years on record. The same ranking also applies to NOAA/NCEI data (Table 1).

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Table 1. Top five warmest years in the ocean since 1955 for both IAP (updated from Cheng et al., 2017 through 2019) and NOAA/NCEI data (updated from Levitus et al., 2012 through 2016). The OHC values are the anomalies (units: ZJ) relative to the 1981–2010 average.

Rank	Year	IAP	NOAA/NCEI
1	2019	228	217
2	2018	203	196
3	2017	193	189
4	2015	185	180
5	2016	180	164

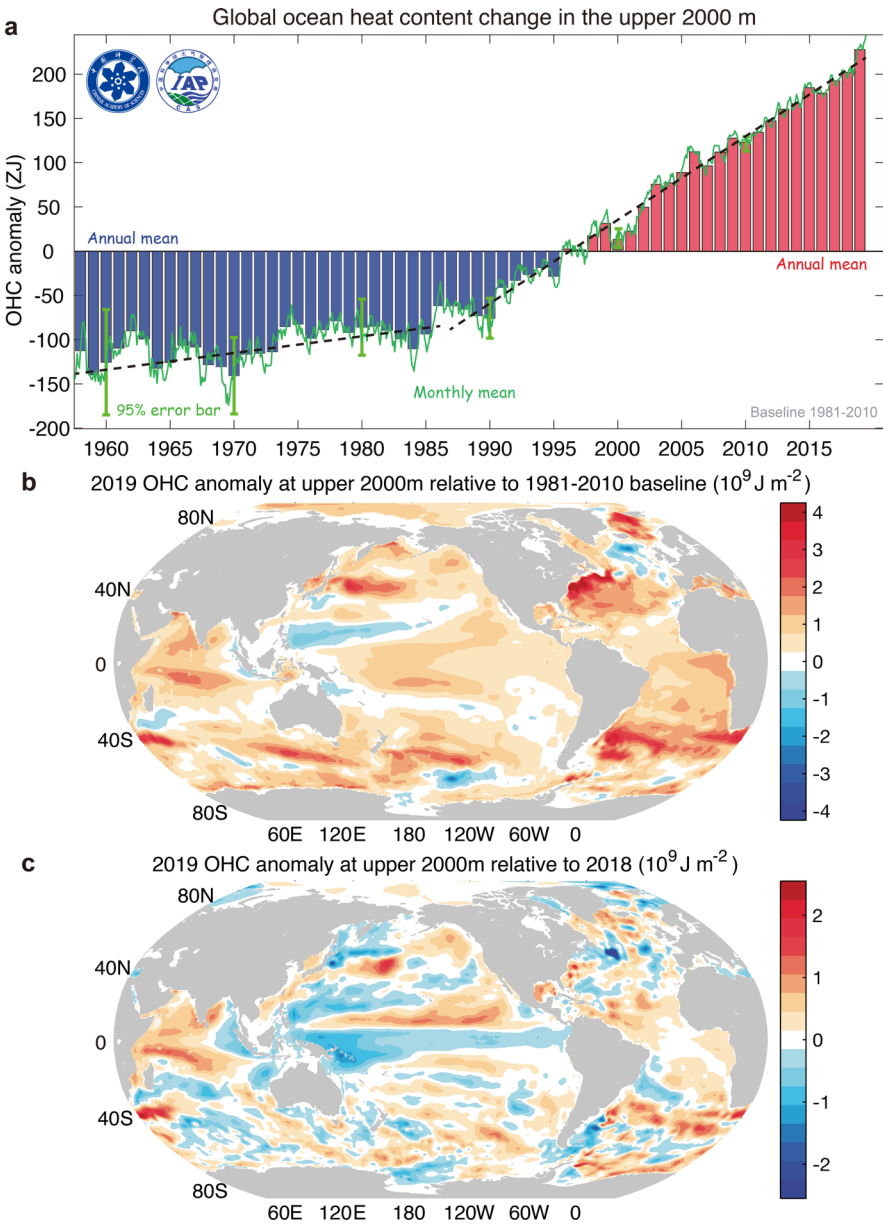


Fig. 1. (a) Upper 2000 m OHC from 1955 through 2019. The histogram represents annual anomalies (units: ZJ), wherein positive anomalies relative to a 1981–2010 baseline are shown as red bars and negative anomalies as blue. The two black dashed lines are the linear trends over 1955–86 and 1987–2019, respectively. (b) Ocean heat content anomaly in 2019 relative to 1981–2010 baseline. (c) difference of OHC at upper 2000 m between 2019 and 2018. Units: 10^9 J m^{-2} . Source: IAP data.

Since a single annual OHC value can be impacted by internal variabilities (i.e., El Niño–Southern Oscillation, ENSO) or instrumental errors, long-term trends are much more important than any individual year for showing climate change; and such trends have been calculated in this work. Simple least-squares trend lines were created for two approximately equal-duration time periods (period 1 = 1955–86; period 2 = 1987–2019). The warming rates were obtained by the rate of change of the data in Fig. 1. For the first time period, the warming was relatively constant at $\sim 2.1 \pm 0.5 \text{ ZJ yr}^{-1}$. However, the more recent warming was $\sim 450\%$ that of the earlier warming ($9.4 \pm 0.2 \text{ ZJ yr}^{-1}$, equal to 0.58 W m^{-2} averaged over the Earth's surface), reflecting a major increase in the rate of global climate change (Cheng et al., 2019a). Note that the selection of end-points, the use of smoothing, and other factors may lead to slightly different rates of warming than those reported here, but in no way change the central findings of this report.

Based on the updated OHC data throughout 2019 from the IAP, a revised ocean energy budget from 0 m to ocean bottom is provided (Fig. 2), along with the OHC below 2000 m adapted from Purkey and Johnson (2010) (use a linear increase of $1.15 \pm 1.0 \text{ ZJ yr}^{-1}$ after 1992). The deep OHC change below 2000 m was extended to 1960 by assuming a zero heating rate before 1991, consistent with Rhein et al., (2013) and Cheng et al., (2017). The new results indicate a total full-depth ocean warming of $370 \pm 81 \text{ ZJ}$ (equal to a net heating of $0.38 \pm 0.08 \text{ W m}^{-2}$ over the global surface) from 1960 to 2019, with contributions of 41.0%, 21.5%, 28.6% and 8.9% from the 0–300-m, 300–700-m, 700–2000-m, and below-2000-m layers, respectively. The uncertainty in the OHC record has been greatly reduced in recent decades through improvements in the ocean observation system (Meyssignac et al., 2019), which is represented by error bars in Fig. 1 (The detailed calculation of the uncertainty range can be found in Cheng et al., 2017). Furthermore, the penetration of heat into the deep oceans is evident (Fig. 2, 3), primarily through Atlantic Ocean and Southern Ocean (Fig. 3).

The heating was distributed throughout the world's oceans, with the vast majority of regions showing an increase in thermal energy (Figs. 1b, 3). The Atlantic Ocean and Southern Ocean (especially near the Antarctic Circumpolar Current, 40° – 60°S) continued to show a larger warming compared to most of the other basins (Cheng et al., 2019c). It is well-established that the Southern Ocean has taken up most of the global warming heat since 1970: multiple observations show that the ocean south of 30°S at 0–2000 m was responsible for 35%–43% of the increase in global OHC within 1970–2017 (Meredith et al., 2019). Also, this is associated with a net northward heat transport by the oceans across the equator (Trenberth and Zhang, 2019). Such meridional oceanic heat transport has influenced recent marine heatwaves in the Tas-

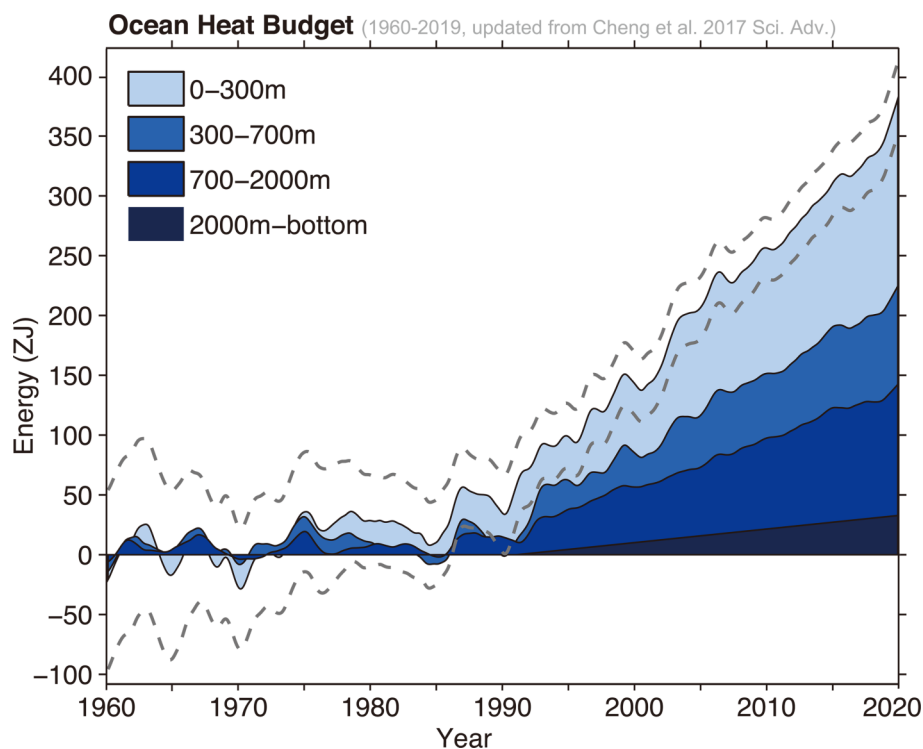


Fig. 2. Ocean heat budget from 1960 to 2019 based on IAP analysis data from 0 to 2000 m, and from Purkey and Johnson (2010) for deep ocean change below 2000 m (units: ZJ). Figure updated from Cheng et al. (2017). The anomalies are related to 1958–1962 baseline, and the time series are smoothed by LOWESS (locally weighted scatterplot smoothing) with span width of 24 months. The gray dashed lines are the 95% confidence interval of the total ocean heat budget.

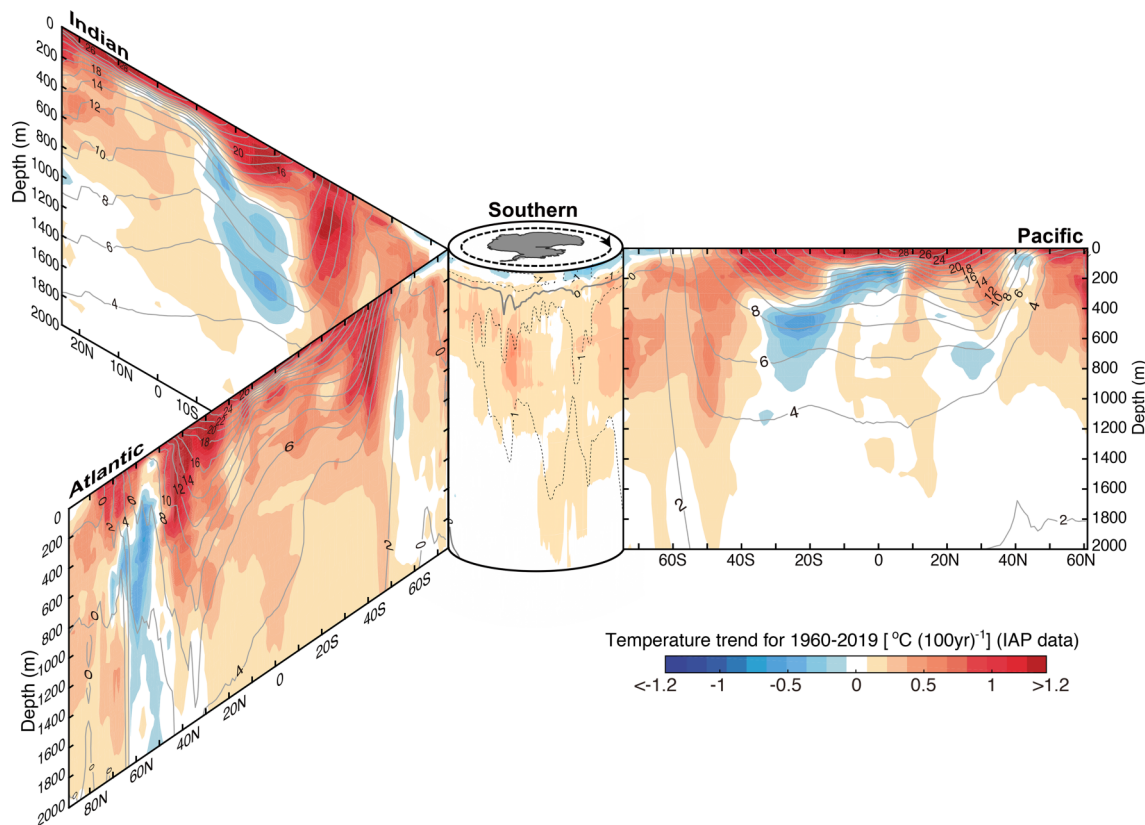


Fig. 3. Vertical section of the ocean temperature trends within 1960 to 2019 from sea surface to 2000 m (60-year Ordinary Least Squares linear trend). Shown are the zonal mean sections in each ocean basin organized around the Southern Ocean (south of 60°S) in the center. Black contours show the associated climatological mean temperature with intervals of 2°C (In the Southern Ocean, 1°C intervals are provided in dashed contours). IAP gridded data are used.

man Sea, which imparts a profound influence on marine fisheries and ecosystems (Behrens et al., 2019; Trenberth and Zhang, 2019). Interestingly, many of the major marine heatwave events in recent years (Collins et al., 2019) have been located near the strong ocean heating regions, such as the Mediterranean Sea, North Pacific (nicknamed “The Blob”), equatorial central Pacific, Tasman Sea, and North Atlantic ($> 40^\circ\text{N}$) (Fig. 1b, 3). Marine heatwaves and other stresses caused by ocean environmental changes clearly pose high risks to biodiversity and fisheries, and cause economic losses.

In the tropics, the spatial pattern of OHC anomalies in 2019 relative to 2018 (Fig. 1c) also reveals the impact of ENSO, which is the dominant mode of the Earth/ocean’s energy variability on interannual scales (Trenberth et al., 2002; Roemmich and Gilson, 2011; Cheng et al., 2019b). ENSO is associated with both heat redistributions within the ocean and heat exchange between the oceans and atmosphere. From late 2018 to early 2019, El Niño formed, and after the summer of 2019 the El Niño event decayed and a neutral phase persisted throughout the second half of the year. From La Niña (early 2018) to El Niño (later 2019), heat discharged from the equatorial and South Pacific Ocean (20°S – 5°N) into the Northeast Pacific Ocean (5° – 20°N), resulting in a cooling of the Pacific Ocean within 20°S – 5°N and a warming of the Pacific/Indian Ocean within 5° – 20°N on the OHC map in Fig. 1c.

The ocean heating is irrefutable, and a key measure of the Earth’s energy imbalance: the excess GHGs in the air trap more heat inside the climate system and drives global warming. More than 90% of the heat accumulates in the ocean because of its large heat capacity, and the remaining heating manifests as atmospheric warming, a drying and warming land-mass, and melting of land and sea ice. There are no reasonable alternatives aside from anthropogenic emissions of heat-trapping gases (IPCC, 2001, 2007, 2013, 2019; USGCRP, 2017). Increased ocean temperatures lead to rising sea levels (thermal expansion of the ocean and added mass from melting land ice) (Oppenheimer et al., 2019). According to the altimetry satellite record, the past 10 years are also the highest in global mean sea level since 1900 (IPCC, 2019).

Increases in ocean temperature reduce dissolved oxygen in the ocean and significantly affect sea life, particularly corals and other temperature- and chemistry-sensitive organisms (Abram et al., 2019; Bindoff et al., 2019). The increasing heat increases evaporation, and the extra moisture in the warmer atmosphere nourishes heavy rains and promotes flooding (Trenberth et al., 2003; Held and Soden, 2006; Trenberth, 2011; Collins et al., 2019), leading to a more extreme hydrological cycle and more extreme weather (in particular hurricanes and typhoons) (Trenberth et al., 2018). It is one of the key reas-

ons why the Earth has experienced increasing catastrophic fires in the Amazon, California, and Australia in 2019 (extending into 2020 for Australia).

It is important to note that ocean warming will continue even if the global mean surface air temperature can be stabilized at or below 2°C (the key policy target of the Paris Agreement) in the 21st century (Cheng et al., 2019a; IPCC, 2019), due to the long-term commitment of ocean changes driven by GHGs. Here, the term “commitment” means that the ocean (and some other components in the Earth system, such as the large ice sheets) are slow to respond and equilibrate, and will continue to change even after radiative forcing stabilizes (Abram et al., 2019). However, the rates and magnitudes of ocean warming and the associated risks will be smaller with lower GHG emissions (Cheng et al., 2019a; IPCC, 2019). Hence, the rate of increase can be reduced by appropriate human actions that lead to rapid reductions in GHG emissions (Cheng et al., 2019a; IPCC, 2019), thereby reducing the risks to humans and other life on Earth.

Data and methods

The method used here, developed at the IAP, involves calculating ocean temperatures down to 2000 m using all available observations from various measurement devices from the World Ocean Database (WOD) of the National Oceanic and Atmospheric Administration/National Center for Environmental Information (NOAA/NCEI). The data are available at <http://159.226.119.60/cheng/> and www.mecp.org.cn/. The primary input data are bias-corrected XBT and other available measurements (i.e. Argo, moorings, gliders, CTD, bottle, MBT etc.) from WOD. Model simulations were used to guide the gap-filling method from point measurements to the grid, while sampling error was estimated by sub-sampling the Argo data at the locations of the earlier observations (a full description of the method can be found in Cheng et al., 2017). Since approximately 2005, the Argo observing network was instituted, which significantly improved the ocean-measurement capability (Argo, 2000). The Argo data were collected and made freely available by the International Argo Program and the national programs that contribute to it (<http://www.argo.ucsd.edu>; <http://argo.jcommops.org>). The Argo Program is part of the Global Ocean Observing System. Altimetry sea level information comes from <https://www.aviso.altimetry.fr/en/data/products/ocean-indicators-products/mean-sea-level.html>. ENSO information can be found at https://origin.cpc.ncep.noaa.gov/products/analysis_monitoring/ensostuff/ONI_v5.php. NOAA/NCEI ocean heat content data can be found in https://www.nodc.noaa.gov/OC5/3M_HEAT_CONTENT/.

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REFERENCES

- Abram, N., and Coauthors, 2019: Framing and context of the report. *IPCC Special Report on the Ocean and Cryosphere in a Changing Climate*, H.-O. Pörtner et al., Eds. (in press)
- Behrens, E., D. Fernandez, and P. Sutton, 2019: Meridional oceanic heat transport influences marine heatwaves in the tasman sea on inter-annual to decadal timescales. *Frontiers in Marine Science*, **6**, 228, <https://doi.org/10.3389/fmars.2019.00228>.
- Bindoff, N. L., and Coauthors, 2019: Changing ocean, marine ecosystems, and dependent communities. *IPCC Special Report on the Ocean and Cryosphere in a Changing Climate*, H.-O. Pörtner et al., Eds. (in press).
- Cheng, L. J., and J. Zhu, 2018: 2017 was the warmest year on record for the global ocean. *Adv. Atmos. Sci.*, **35**, 261–263, <https://doi.org/10.1007/s00376-018-8011-z>.
- Cheng, L. J., K. E. Trenberth, J. Fasullo, T. Boyer, J. Abraham, and J. Zhu, 2017: Improved estimates of ocean heat content from 1960 to 2015. *Science Advances*, **3**, e1601545, <https://doi.org/10.1126/sciadv.1601545>.
- Cheng, L. J., K. E. Trenberth, J. Fasullo, J. Abraham, T. P. Boyer, K. Von Schuckmann, and J. Zhu, 2018: Taking the pulse of the planet. *Eos*, **99**, 14–16, <https://doi.org/10.1029/2017EO081839>.
- Cheng, L. J., J. Abraham, Z. Hausfather, and K. E. Trenberth, 2019a: How fast are the oceans warming? *Science*, **363**, 128–129, <https://doi.org/10.1126/science.aav7619>.
- Cheng, L. J., K. E. Trenberth, J. T. Fasullo, M. Mayer, M. A. Balmaseda, and J. Zhu, 2019b: Evolution of ocean heat content related to ENSO. *J. Climate*, **32**, 3529–3556, <https://doi.org/10.1175/JCLI-D-18-0607.1>.
- Cheng, L. J., and Coauthors, 2019c: 2018 continues record global ocean warming. *Adv. Atmos. Sci.*, **36**, 249–252, <https://doi.org/10.1007/s00376-019-8276-x>.
- Collins, M., and Coauthors, 2019: Extremes, abrupt changes and managing risks. *IPCC Special Report on the Ocean and Cryosphere in*

- a *Changing Climate*, H.-O. Pörtner et al., Eds., in press.
- Held, I. M., and B. J. Soden, 2006: Robust responses of the hydrological cycle to global warming. *J. Climate*, **19**, 5686–5699, <https://doi.org/10.1175/jcli3990.1>.
- IPCC, 2001: *Climate Change 2001: the Scientific Basis. Contribution of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change*, J. T. Houghton et al., Eds., Cambridge University Press, 881 pp.
- IPCC, 2007: Climate change 2007: the physical science basis. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change, S. Solomon et al., Eds., *IPCC Fourth Assessment Report. Climate Change 2007, Working Group I Report "The Physical Science Basis"*, Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 996 pp.
- IPCC, 2013: *Climate Change 2013: the Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, T. F. Stocker et al., Eds., Cambridge University Press, 1535 pp.
- IPCC, 2019: Summary for policymakers. *IPCC Special Report on the Ocean and Cryosphere in a Changing Climate*, H.-O. Pörtner et al., Eds. (in press)
- Levitus, S., et al, 2012: World ocean heat content and thermosteric sea level change (0–2000 m), 1955–2010. *Geophys. Res. Lett.*, **39**, L10603, <https://doi.org/10.1029/2012GL051106>.
- Meredith, M., and Coauthors, 2019: Polar regions. *IPCC Special Report on the Ocean and Cryosphere in a Changing Climate*, H.-O. Pörtner et al., Eds. (in press)
- Meyssignac, B., and Coauthors, 2019: Measuring global ocean heat content to estimate the earth energy imbalance. *Frontiers in Marine Science*, **6**, 432, <https://doi.org/10.3389/fmars.2019.00432>.
- Oppenheimer, M., and Coauthors, 2019: Sea level rise and implications for low lying islands, coasts and communities. *IPCC Special Report on the Ocean and Cryosphere in a Changing Climate*, H.-O. Pörtner et al., Eds. (in press)
- Purkey, S. G., and G. C. Johnson, 2010: Warming of global abyssal and deep southern ocean waters between the 1990s and 2000s: Contributions to global heat and sea level rise budgets. *J. Climate*, **23**, 6336–6351, <https://doi.org/10.1175/2010JCLI3682.1>.
- Rhein, M., and Coauthors, 2013: Observations: Ocean. *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, T. F. Stocker et al., Eds., Cambridge University Press, Cambridge, UK, New York, NY, USA.
- Roemmich, D., and J. Gilson, 2011: The global ocean imprint of ENSO. *Geophys. Res. Lett.*, **38**, L13606, <https://doi.org/10.1029/2011GL047992>.
- Trenberth, K. E., 2011: Changes in precipitation with climate change. *Climate Research*, **47**, 123–138, <https://doi.org/10.3354/cr00953>.
- Trenberth, K. E., and Y. X. Zhang, 2019: Observed interhemispheric meridional heat transports and the role of the Indonesian through-flow in the Pacific Ocean. *J. Climate*, **32**, 8523–8536, <https://doi.org/10.1175/JCLI-D-19-0465.1>.
- Trenberth, K. E., J. M. Caron, D. P. Stepaniak, and S. Worley, 2002: Evolution of El Niño-southern Oscillation and global atmospheric surface temperatures. *J. Geophys. Res.*, **107**, 4065, <https://doi.org/10.1029/2000JD000298>.
- Trenberth, K. E., A. G. Dai, R. M. Rasmussen, and D. B. Parsons, 2003: The changing character of precipitation. *Bull. Amer. Meteorol. Soc.*, **84**, 1205–1218, <https://doi.org/10.1175/BAMS-84-9-1205>.
- Trenberth, K. E., L. J. Cheng, P. Jacobs, Y. X. Zhang, and J. Fasullo, 2018: Hurricane Harvey links to ocean heat content and climate change adaptation. *Earth's Future*, **6**, 730–744, <https://doi.org/10.1029/2018EF000825>.
- Trenberth, K., J. T. Fasullo, K. Von Schuckmann, and L. J. Cheng, 2016: Insights into Earth's energy imbalance from multiple sources. *J. Climate*, **29**, 7495–7505, <https://doi.org/10.1175/JCLI-D-16-0339.1>.
- USGCRP, 2017: Climate science special report: fourth national climate assessment, Volume I [Wuebbles, D.J., D.W. Fahey, K.A. Hibbard, D.J. Dokken, B.C. Stewart, and T.K. Maycock (eds.)]. U.S. Global Change Research Program, Washington, DC, USA, 470 pp, <https://doi.org/10.7930/J0J964J6>.
- Von Schuckmann, K., and Coauthors, 2016: An imperative to monitor Earth's energy imbalance. *Nat. Clim. Change*, **6**, 138–144, <https://doi.org/10.1038/nclimate2876>.